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Simulating Injection Induced Earthquakes: The Effects of Rate-and State-dependent Friction and Injection History

Kayla A. Kroll

Keith B. Richards-Dinger, James H. Dieterich, and Joshua A. White



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What Are Earthquake Simulators?

- Computer codes which
 - Solve simplified forms of the physical equations
 - Typically boundary elements
 - Quasi-static with some approximation to inertia
 - Group the resulting time histories of slip across the elements into earthquake events
 - Multiple earthquake cycles

Earthquake Time-dependency: Rate- and State-dependent Friction

$$\tau = \mu(\sigma - p)$$

(1) Modified Coulomb Criterion

$$\mu = \mu_0 + a \ln\left(\frac{V}{V^*}\right) + b \ln\left(\frac{\theta V^*}{D_c}\right)$$

(2) Rate- and State-dependent friction

$$\tau = (\sigma - p) \left[\mu_0 + \boxed{a \ln\left(\frac{V}{V^*}\right)} \boxed{b \ln\left(\frac{\theta V^*}{D_c}\right)} \right]$$

(3) Constitutive Law

Rate-term State-term

μ_0 : Nominal coefficient of friction

V^* : Reference slip rate

V : Earthquake slip rate

θ : State variable

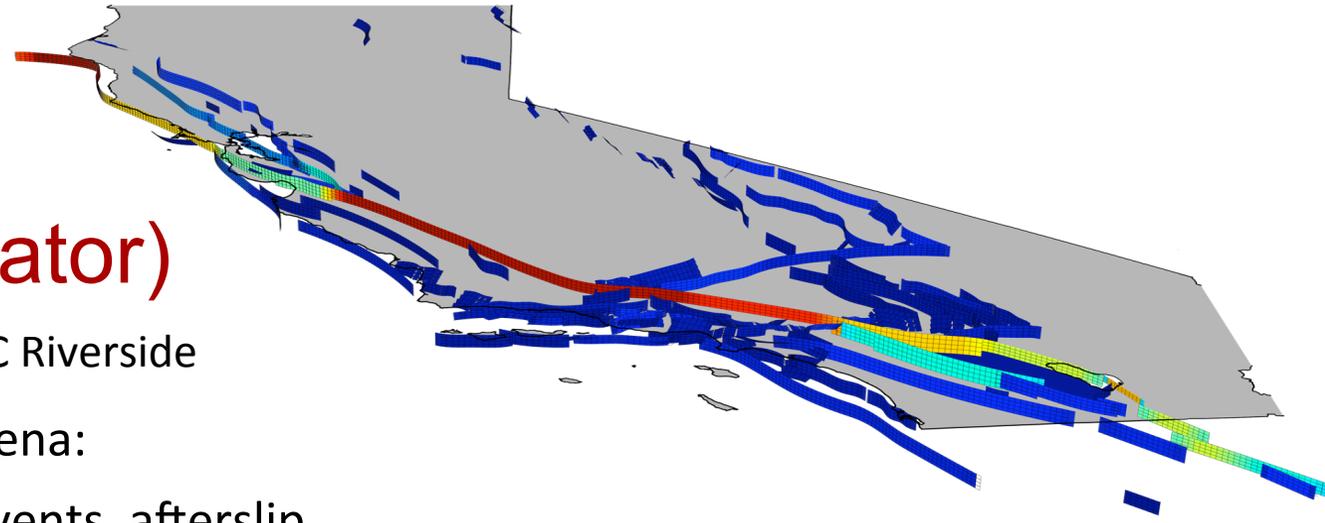
D_c : Characteristic slip distance

a and b : Constitutive parameters describing the material

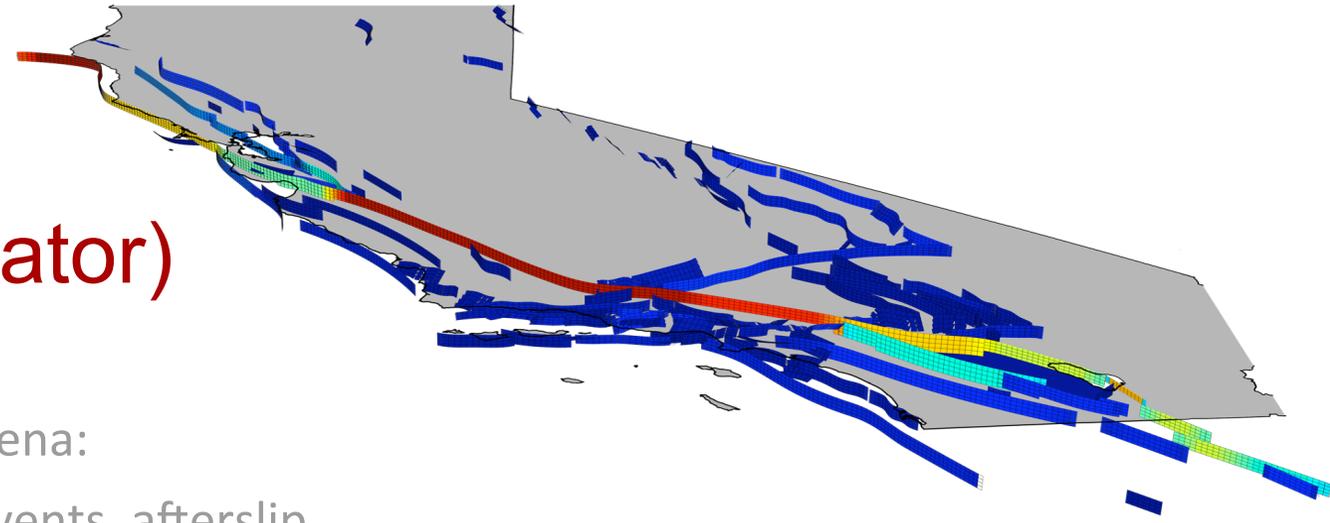
RSQSim (Rate-State earthQuake Simulator)

Developed by Jim Dieterich and Keith Richards-Dinger at UC Riverside

- Comprehensive simulation of fault slip phenomena:
 - earthquakes, continuous creep, slow slip events, afterslip

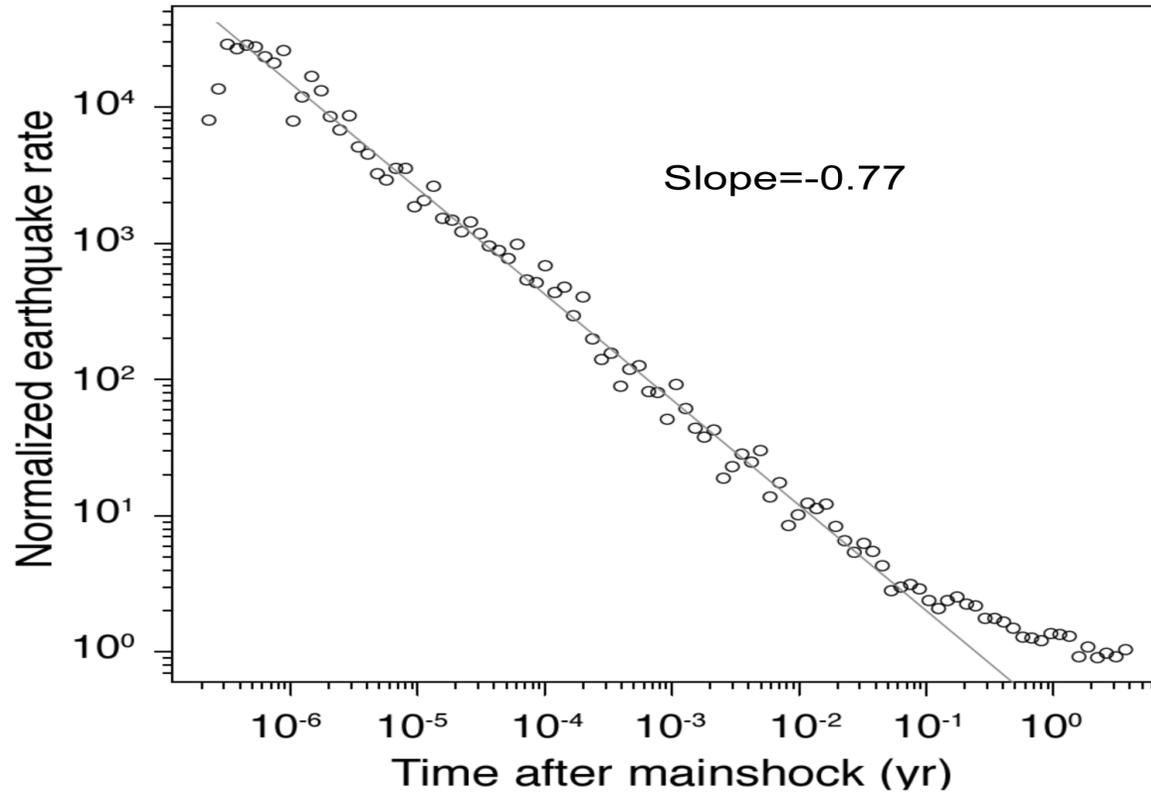
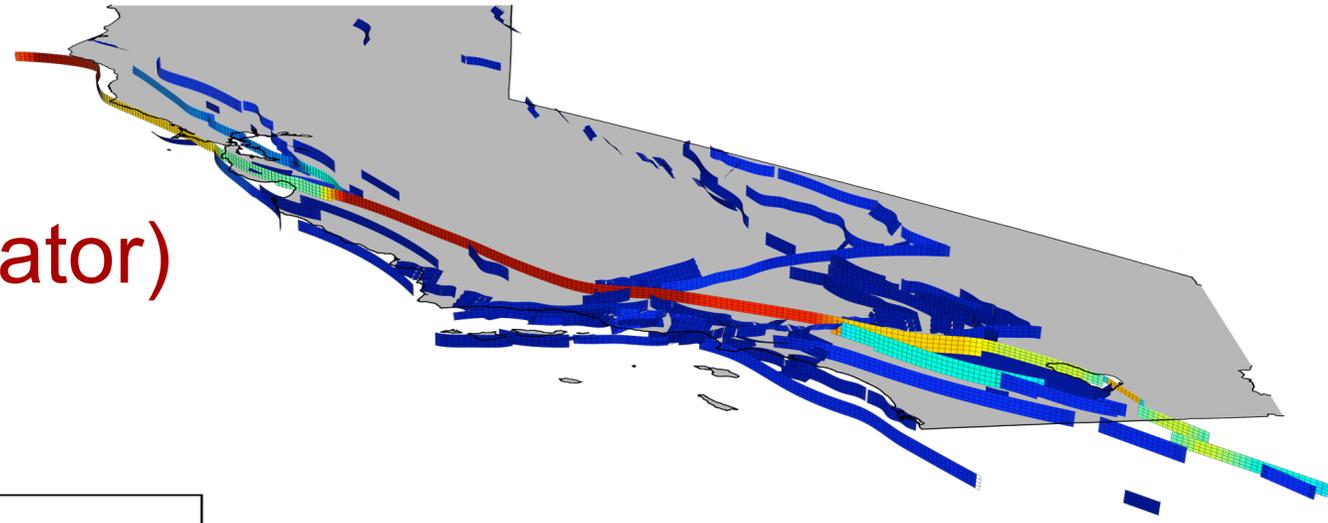


RSQSim (Rate-State earthQuake Simulator)



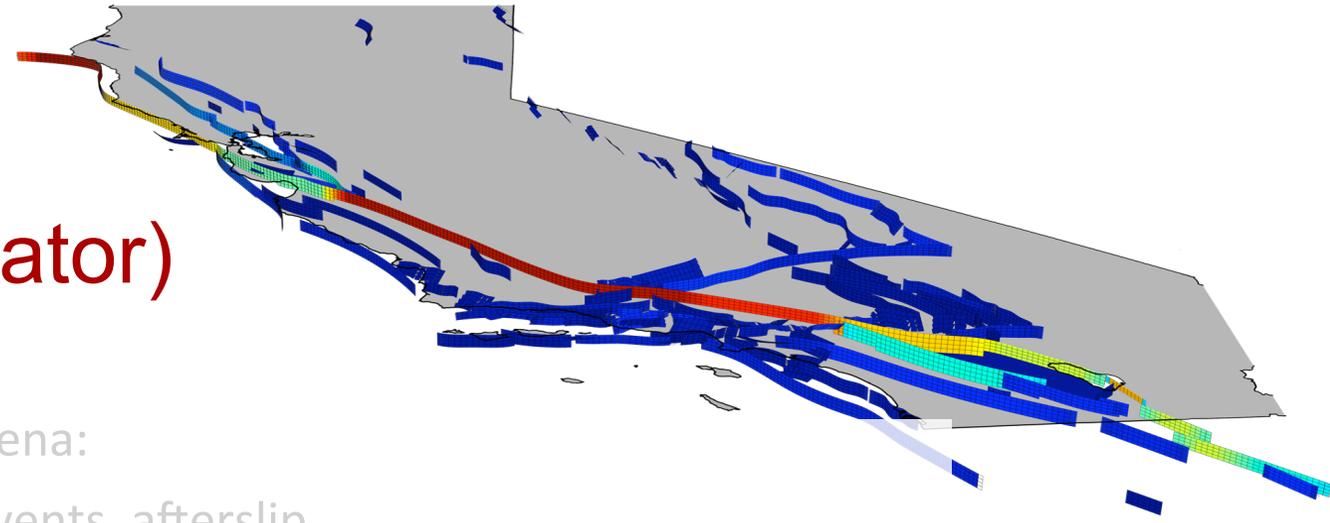
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 - Earthquake clustering effects (aftershocks and foreshocks)

RSQSim (Rate-State earthQuake Simulator)



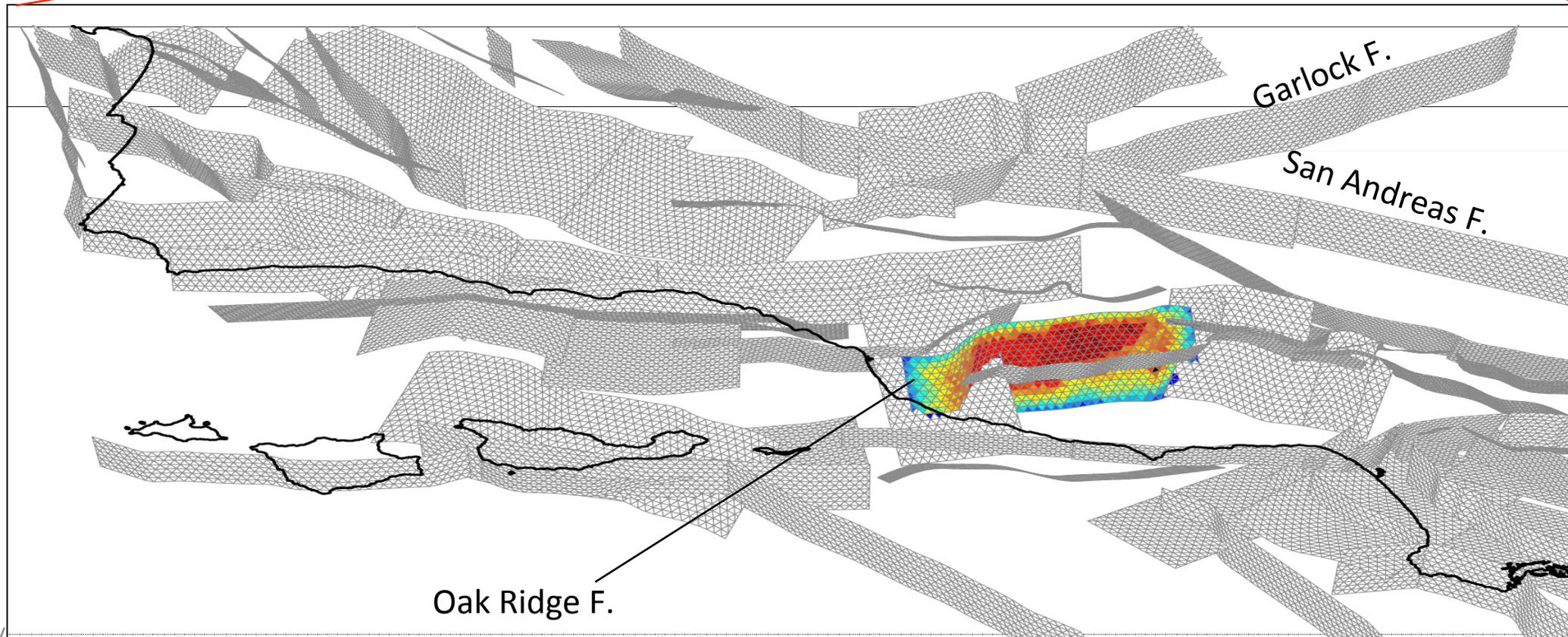
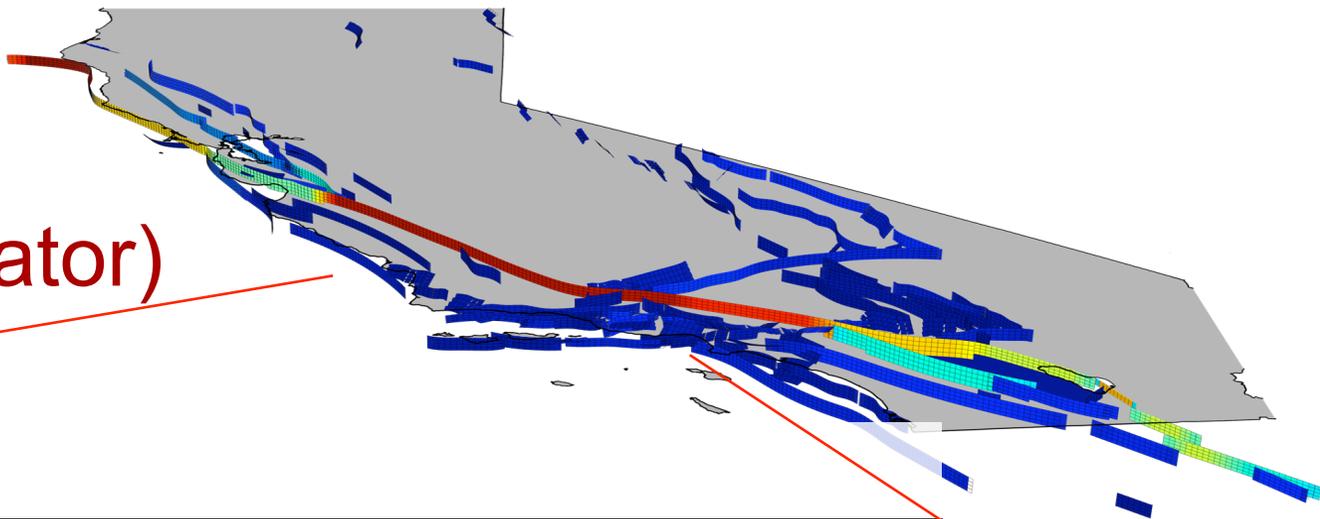
All-California simulation
Aftershocks follow the Omori Law
for aftershock decay with time

RSQSim (Rate-State earthQuake Simulator)

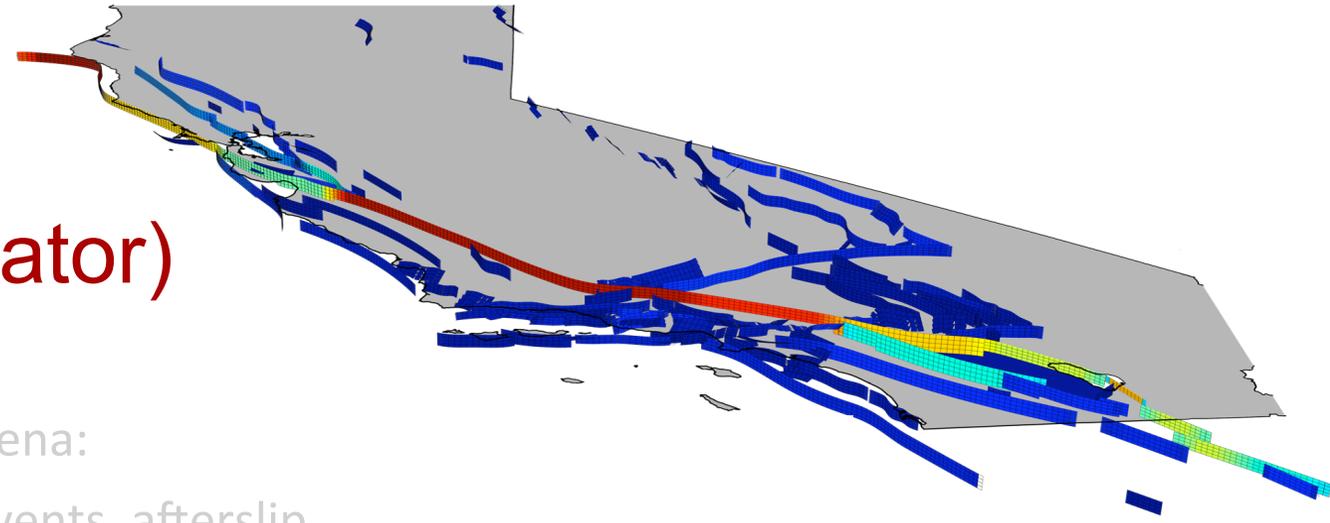


- Comprehensive simulation of fault slip phenomena:
 - earthquakes, continuous creep, slow slip events, afterslip
- Implement rate- and state-dependent friction effects
 - Earthquake clustering effects (aftershocks and foreshocks)
- High resolution models of geometrically complex fault systems
 - Up to 10^6 fault elements
 - Range of earthquake magnitudes $M=3.5$ to $M=8$ (for 1 km^2 triangular elements)

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- High resolution models of geometrically complex fault systems
 - Up to 10^6 fault elements
 - Range of earthquake magnitudes $M=3.5$ to $M=8$ (for 1 km^2 triangular elements)
- Highly efficient code
 - Good statistical characterizations from long simulations of 10^6 earthquakes
 - Repeated simulations to explore parameter space

What do we need to begin a simulation?

- External stressing history (from a reservoir model)
- Fault geometry
- Constitutive parameters
- Tectonic driving stress (perhaps neglect these in relatively aseismic regions)
- Initial stress conditions:
 - *In situ* stress measurements - regional average (from global stress maps)
 - Projection of the regional stress tensor (from global stress maps)
 - Randomly generated heterogeneous field (some fractal distribution)
 - Final stress from a large event is a previous simulation (from RSQSim)
 - Evolved stress from a tectonically driven simulation (from RSQSim)

Modeling Injection-Induced Seismicity with the Physics-Based Earthquake Simulator RSQSim

by James H. Dieterich, Keith B. Richards-Dinger, and Kayla A. Kroll

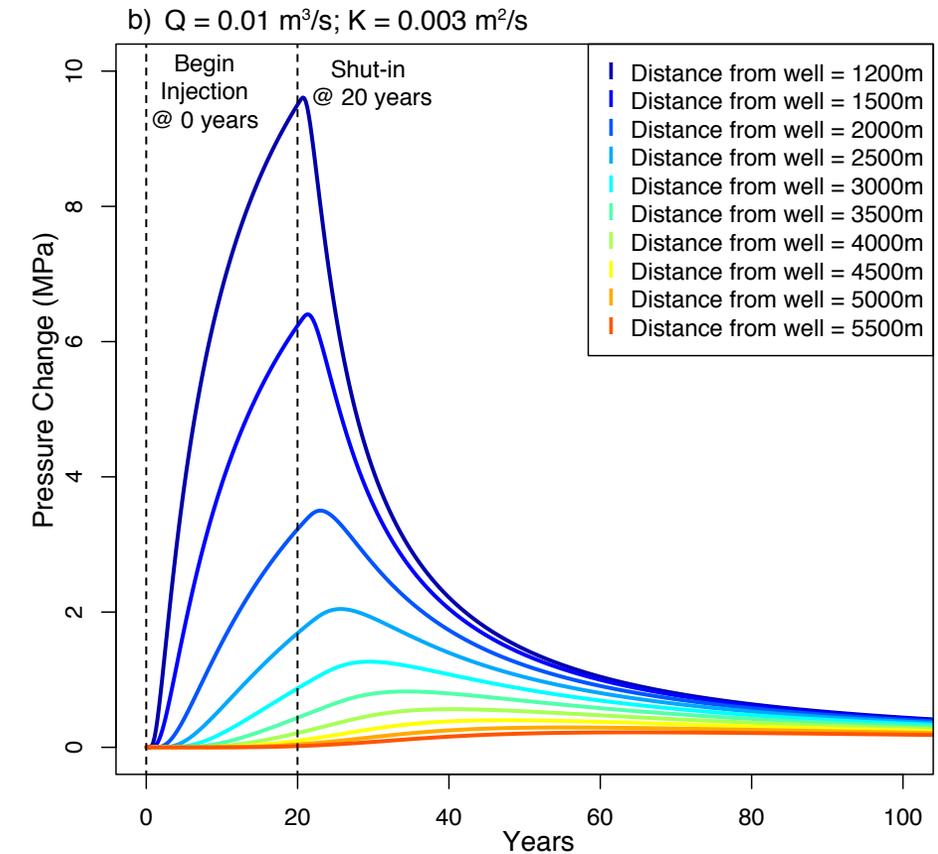
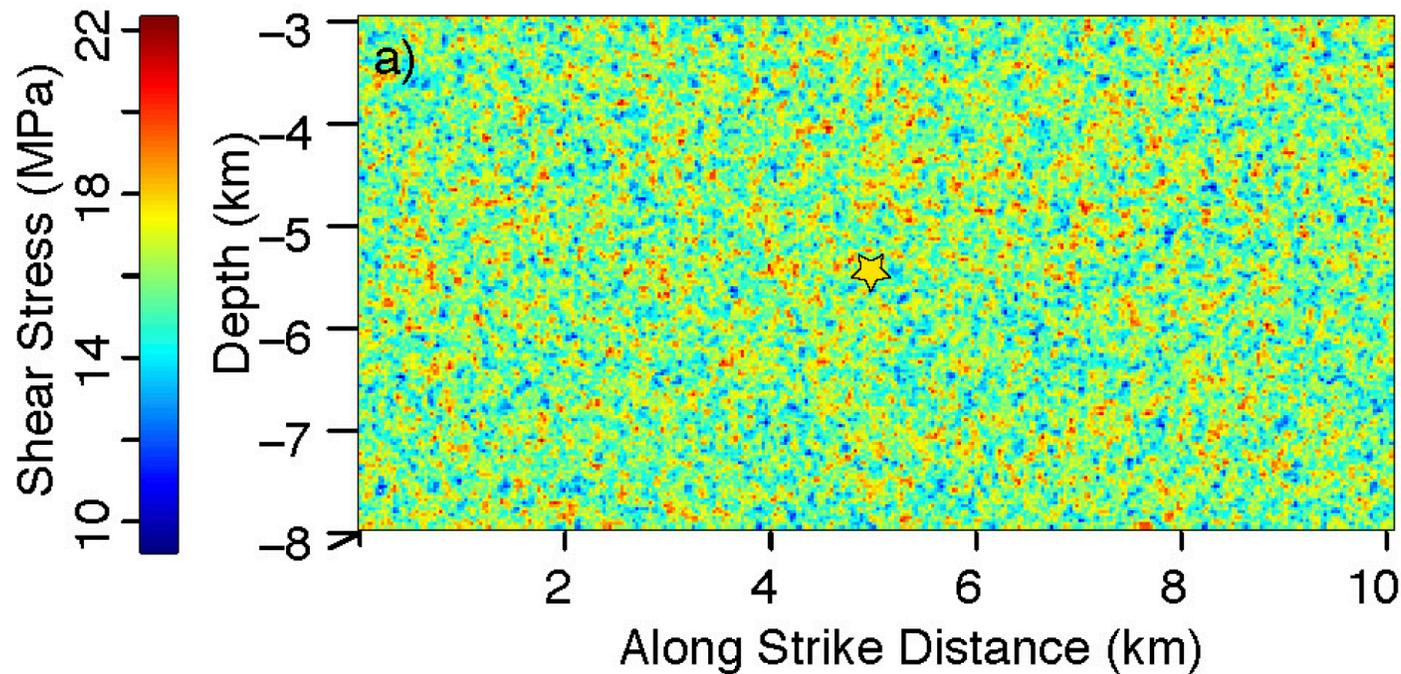
Seismological Research Letters Volume 86, Number 4 July/August 2015 1

Space-time characteristics are highly dependent on the pattern and magnitude the of initial stresses

How do we include effects of fluid injection in RSQSim?

- RSQSim itself knows nothing of pore-fluid pressure diffusion, poroelastic effects, etc.
- Must supply external stressing history
- Geomechanical reservoir model
 - Changes in effective normal stress
 - Poroelastic effects
- Not fully coupled – no feedback
 - seismic slip does not affect the permeability structure, etc.
- Preliminary experiments use a simple analytic expression for pore-fluid diffusion (Wang, 2000).

Fault and “Reservoir” Model



- Linear diffusion model based on analytical solutions for a point-source in an semi-infinite, isotropic half-space (Wang, 2000).
- Variable injection parameters:
 - Well location(s)
 - Injection Rate
 - Diffusivity $\left(\kappa = \frac{k}{\eta\phi c}\right)$; permeability, porosity, compressibility, viscosity

Dependency on Rate-State Parameters

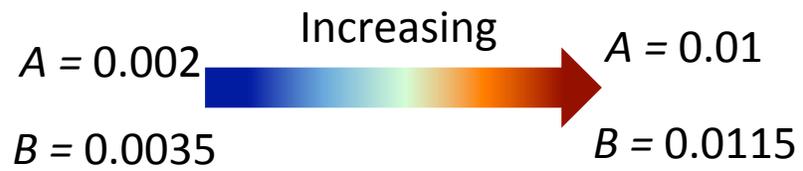
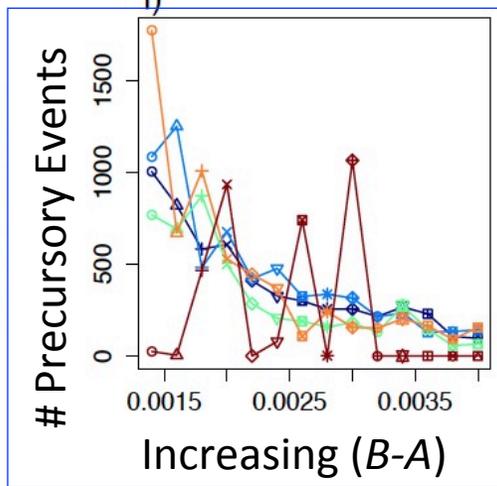
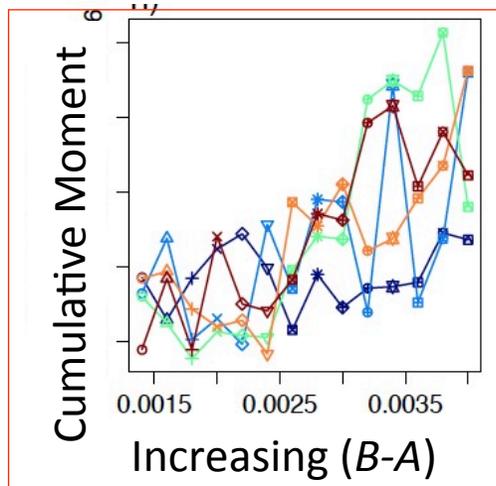
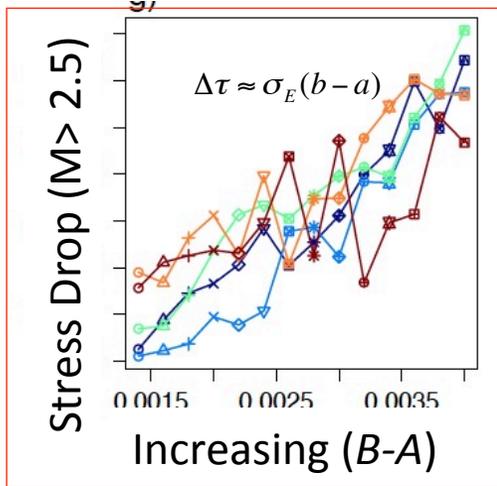
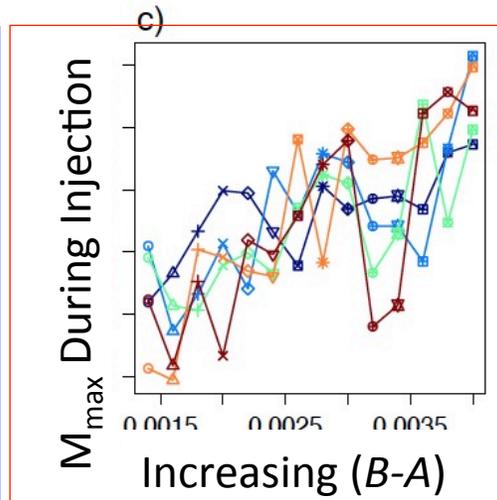
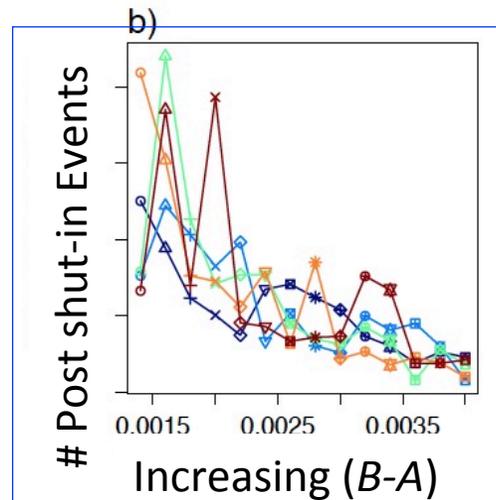
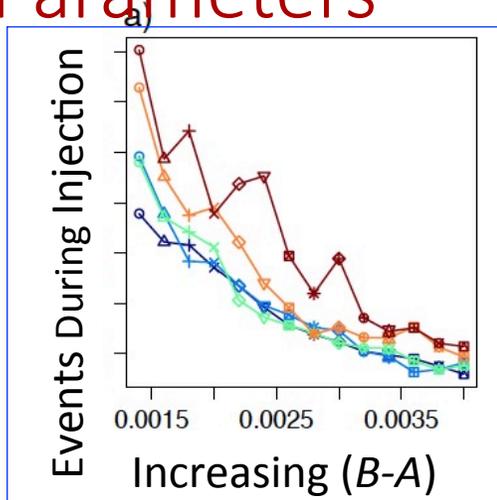
$$\tau = (\sigma - p) \left[\mu_0 \boxed{(b - a)} \ln \left(\frac{\theta V^*}{D_c} \right) \right]$$

Earthquake Stress Drops:

$$\Delta\tau \approx (\sigma - p) \left[(b - a) \ln \left(\frac{\theta V^*}{D_c} \right) \right]$$

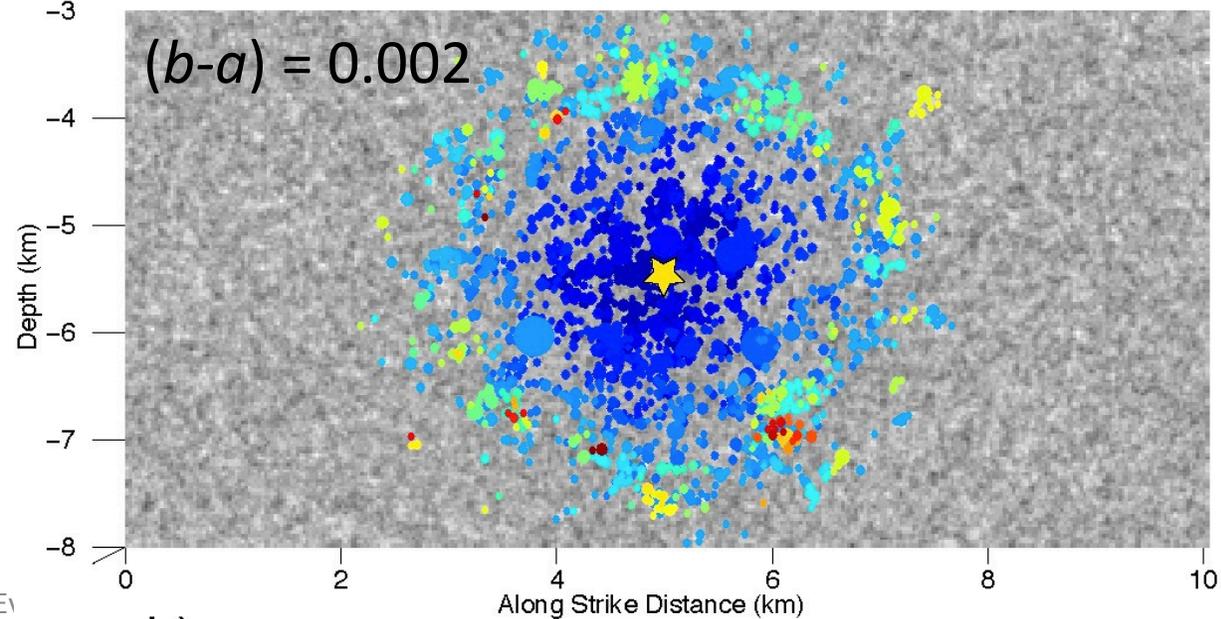
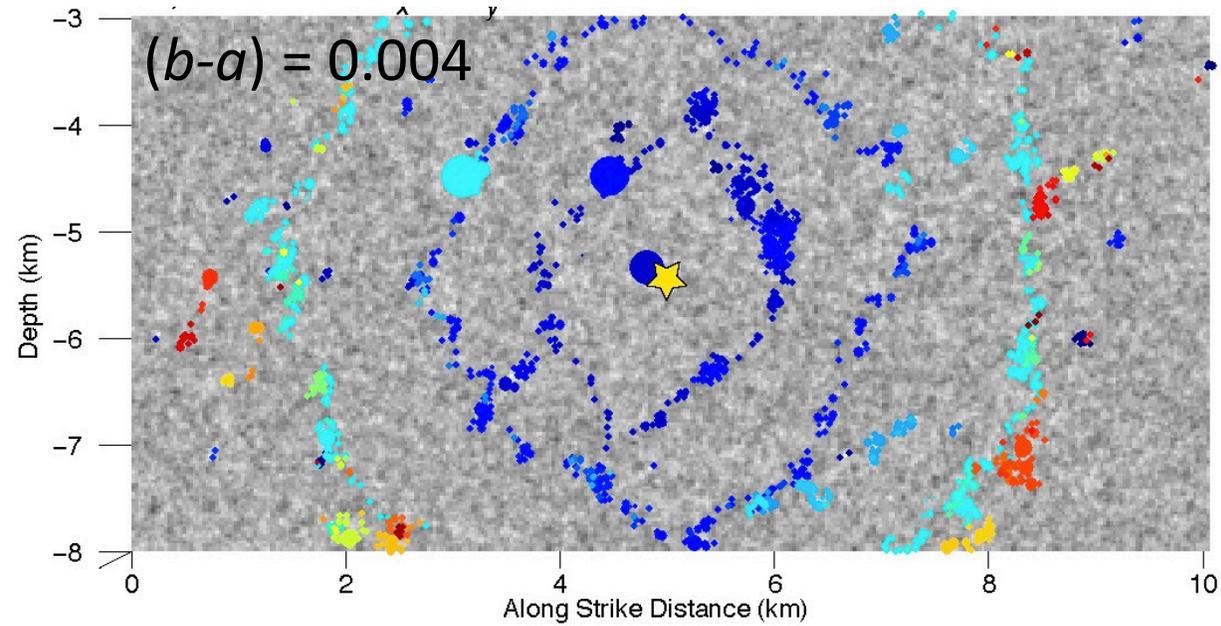
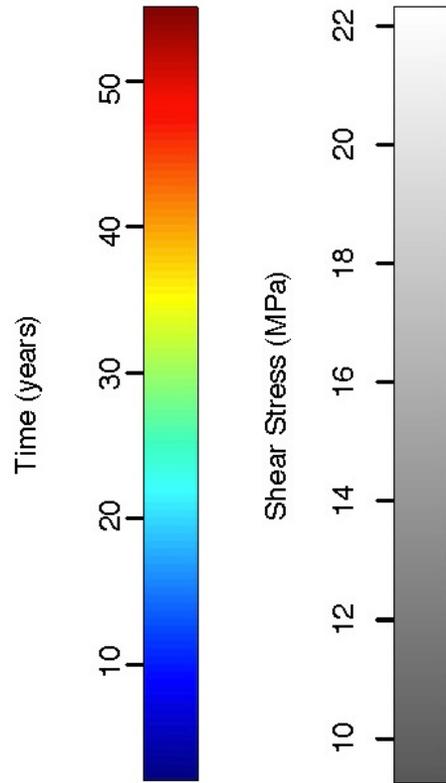
$$\Delta\tau \approx \sigma_E (b - a)$$

$$\Delta\tau = \sigma (\mu_d - \mu_s)$$



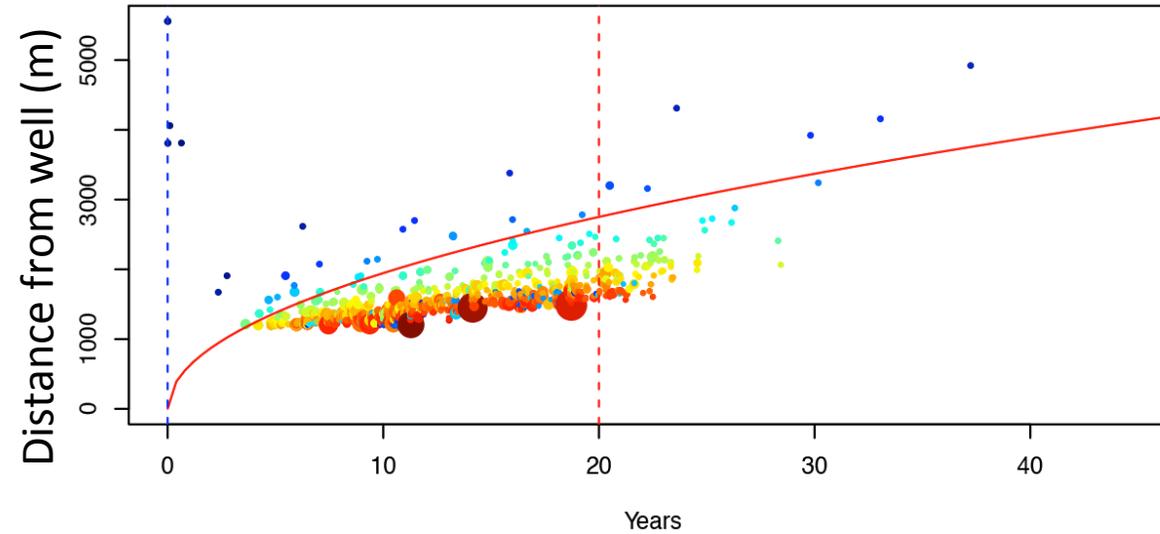
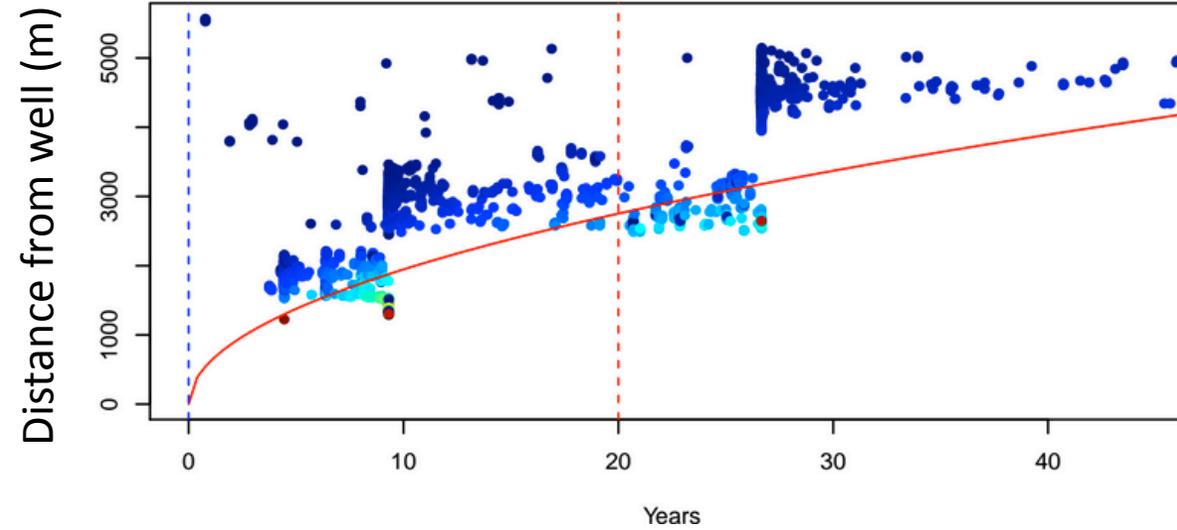
Spatial Variation due to $(b-a)$

** Closest point to well
Is 1200 m

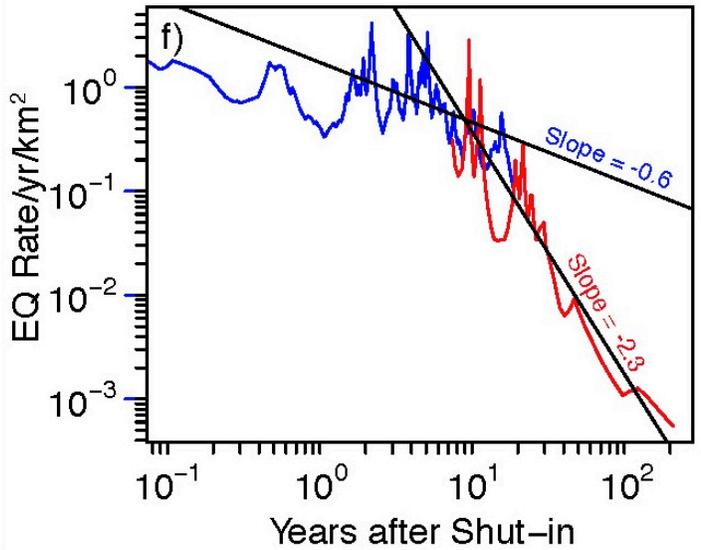
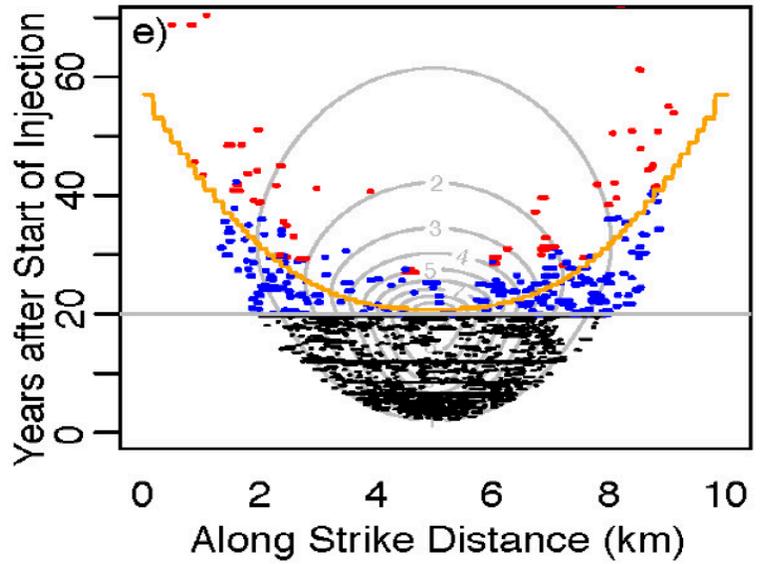
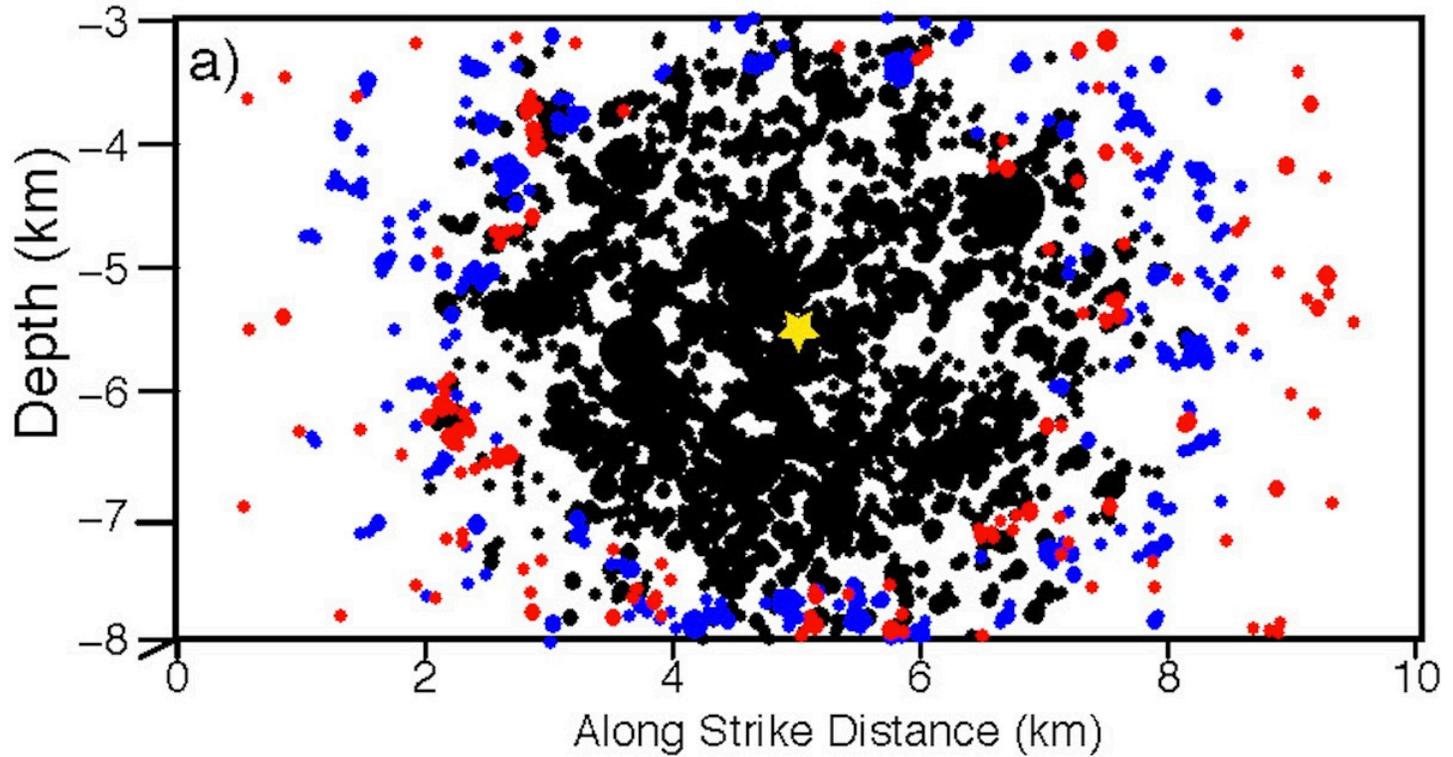


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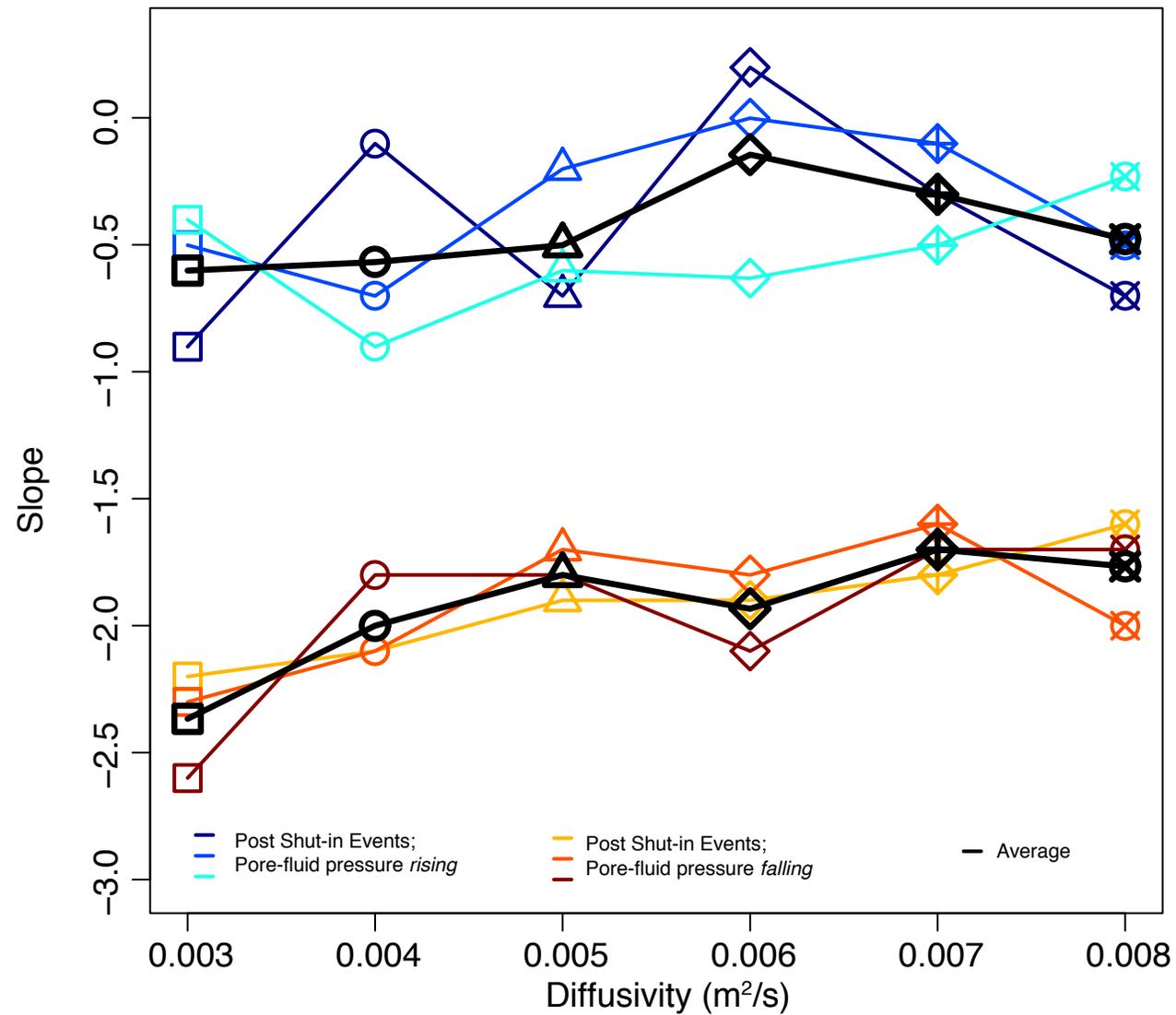


Post Shut-in Seismicity Rates



- Events that occur during the injection period
- Events that occur after shut-in, as the pore-fluid pressure is rising
- Events that occur after shut-in, as the pore-fluid pressure is dropping
- Time at which the pore-fluid pressure starts to decrease (projection from the row of elements at -5500 m depth, even with the well)

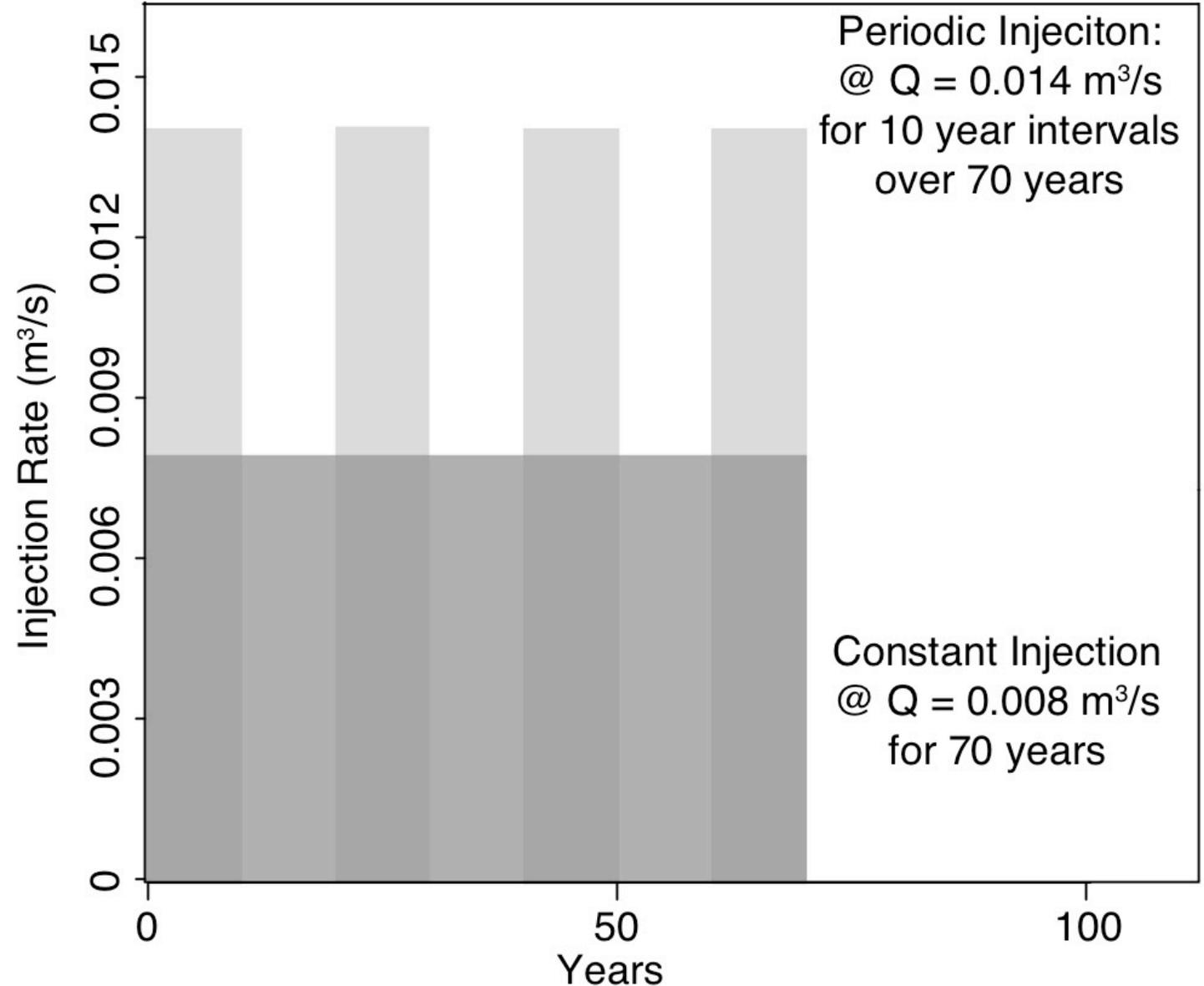
Decay Rate with Diffusivity



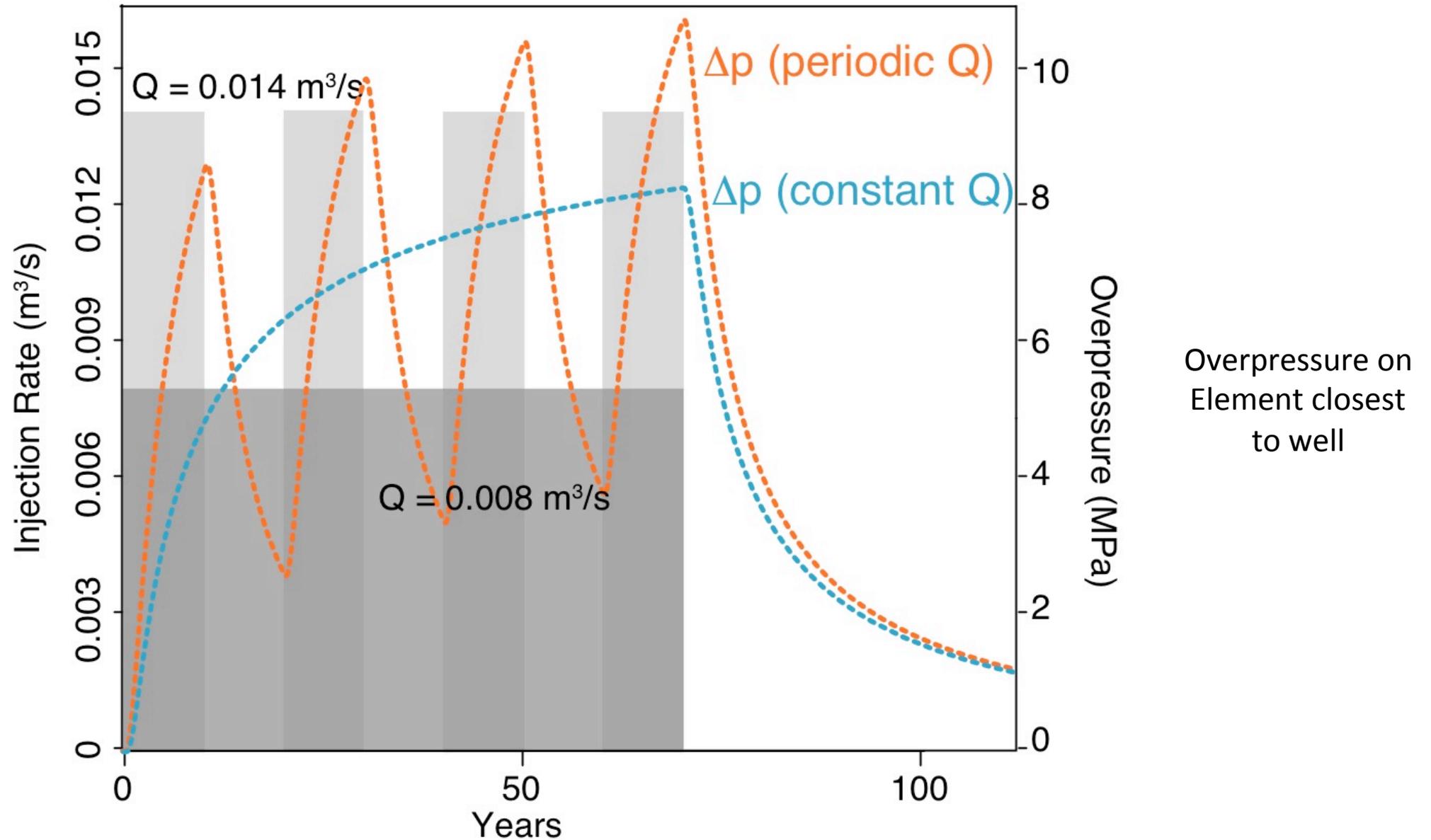
Comparing Injection Histories

Is seismicity controlled by changes in **peak overpressure** or **constant injection rate**?

- Constant injection duration (70 years)
- Constant injected volume ($1.8 \cdot 10^7 \text{ m}^3$)
- Variable injection rate
 - Constant ($0.008 \text{ m}^3/\text{s}$)
 - Periodic ($0.014 \text{ m}^3/\text{s}$)

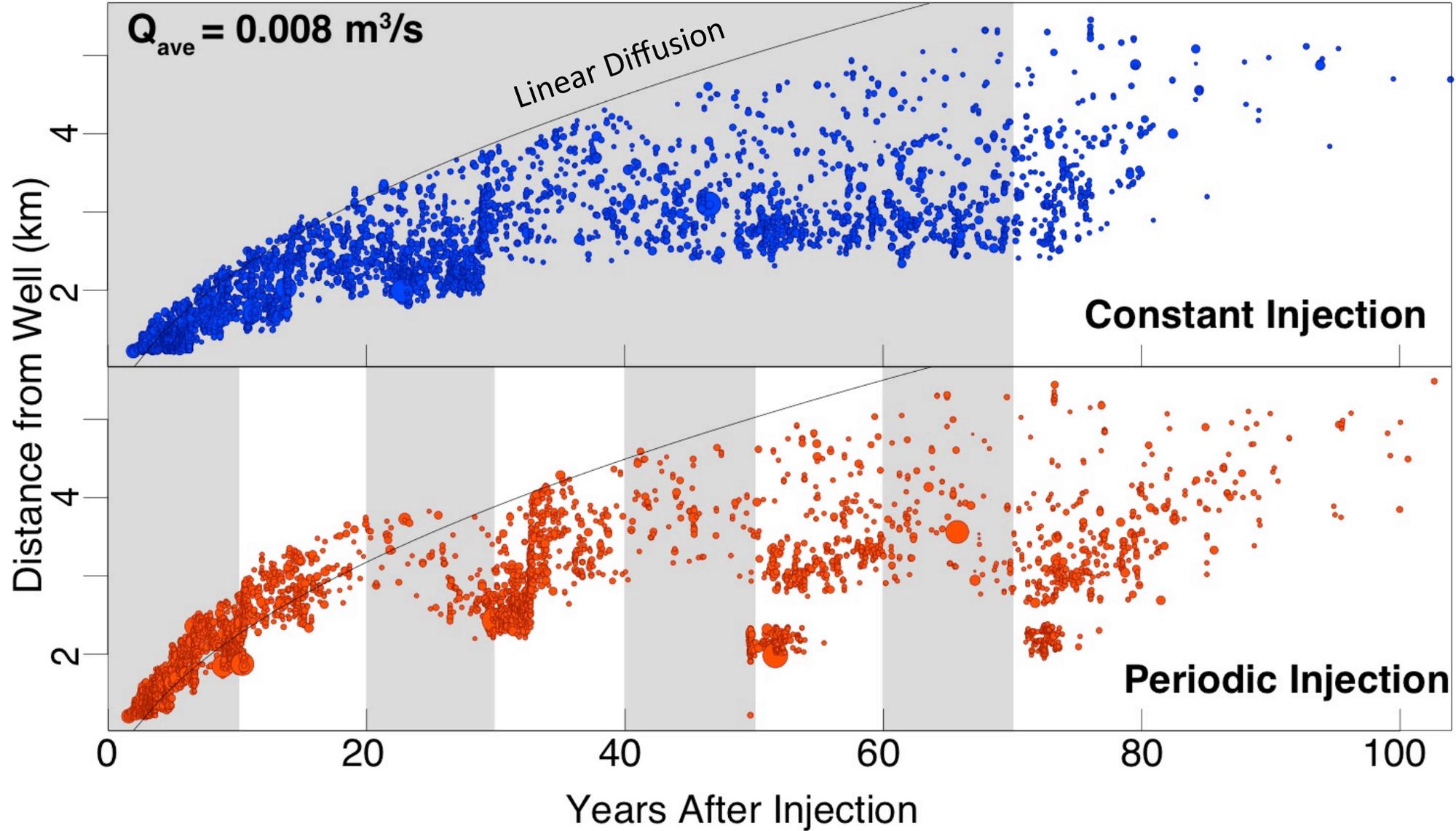


Pressure Change in Response to Constant and Periodic Injection Rate



Results: Spatial Distribution with Time

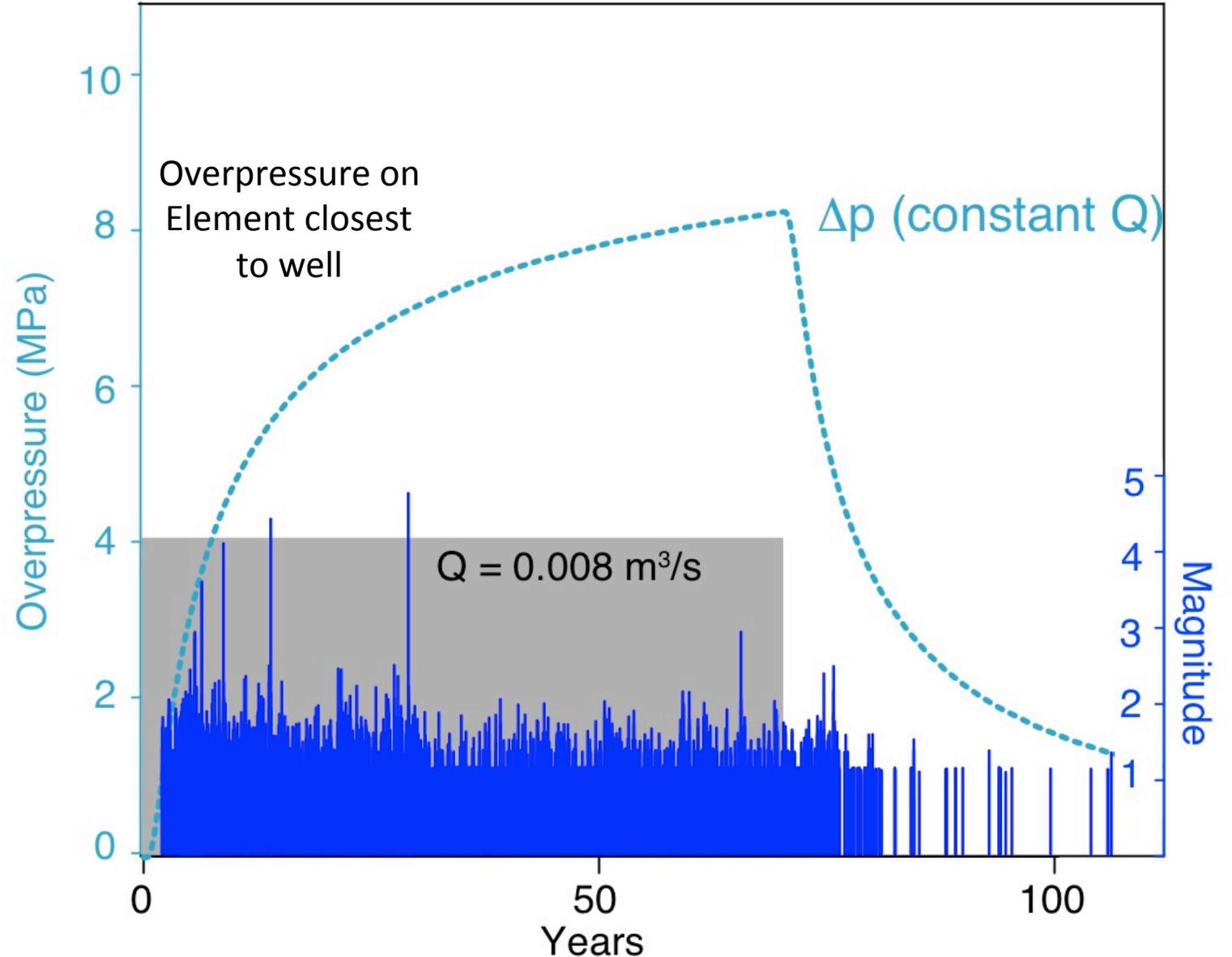
Closest fault element to well = 1.2 km



Scale by earthquake magnitude

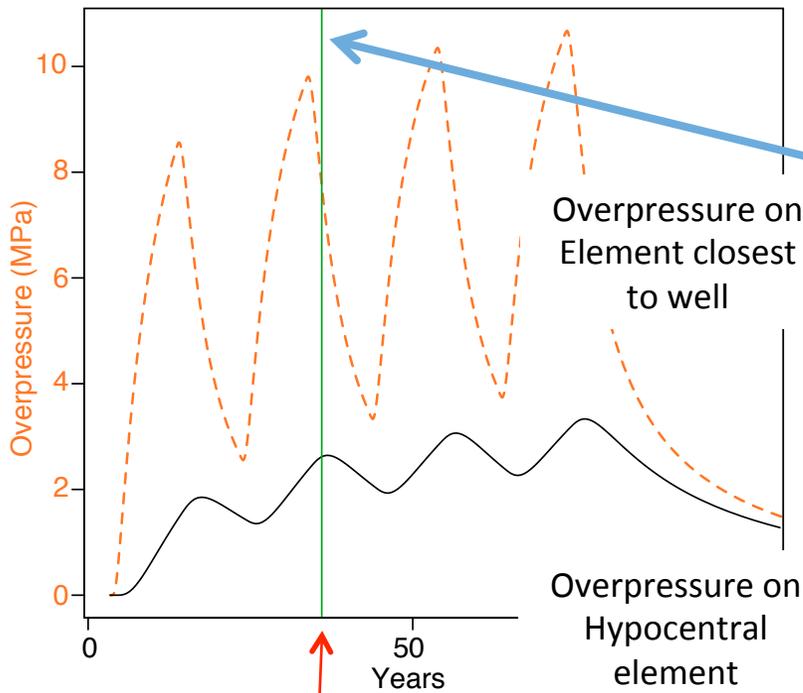
Results: Earthquake Magnitude with Time (Constant Injection Rate)

- Max magnitudes increase with time
- No large, post shut-in events
- 4 $M > 3.5$ events within ~ 20 years
- Relatively constant rate

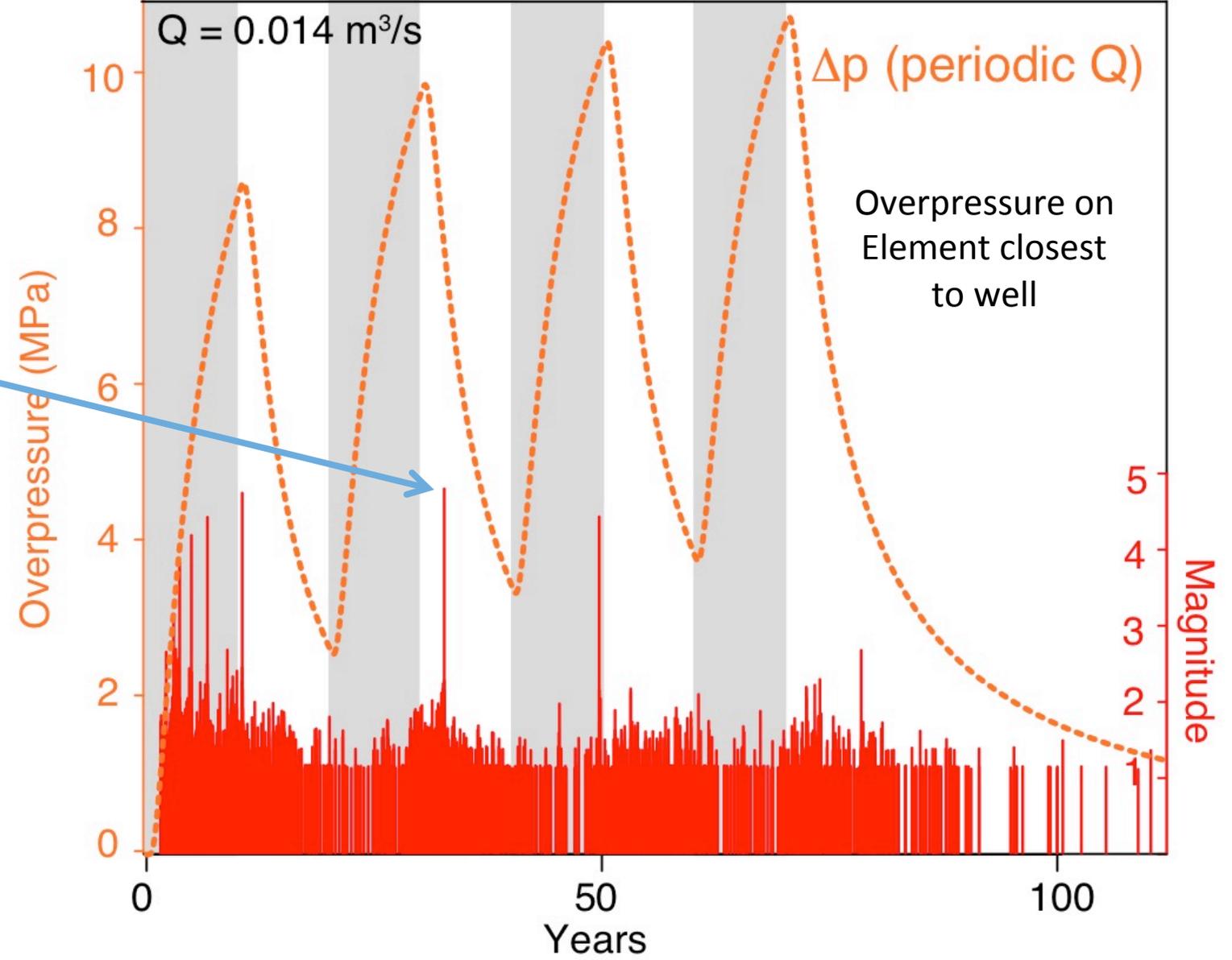


Results: Earthquake Magnitude with Time (Periodic Injection Rate)

- Max magnitudes increase with time
- Most large events occur after shut-in
- Large events occur over ~50 years
- Earthquake rate fluctuates



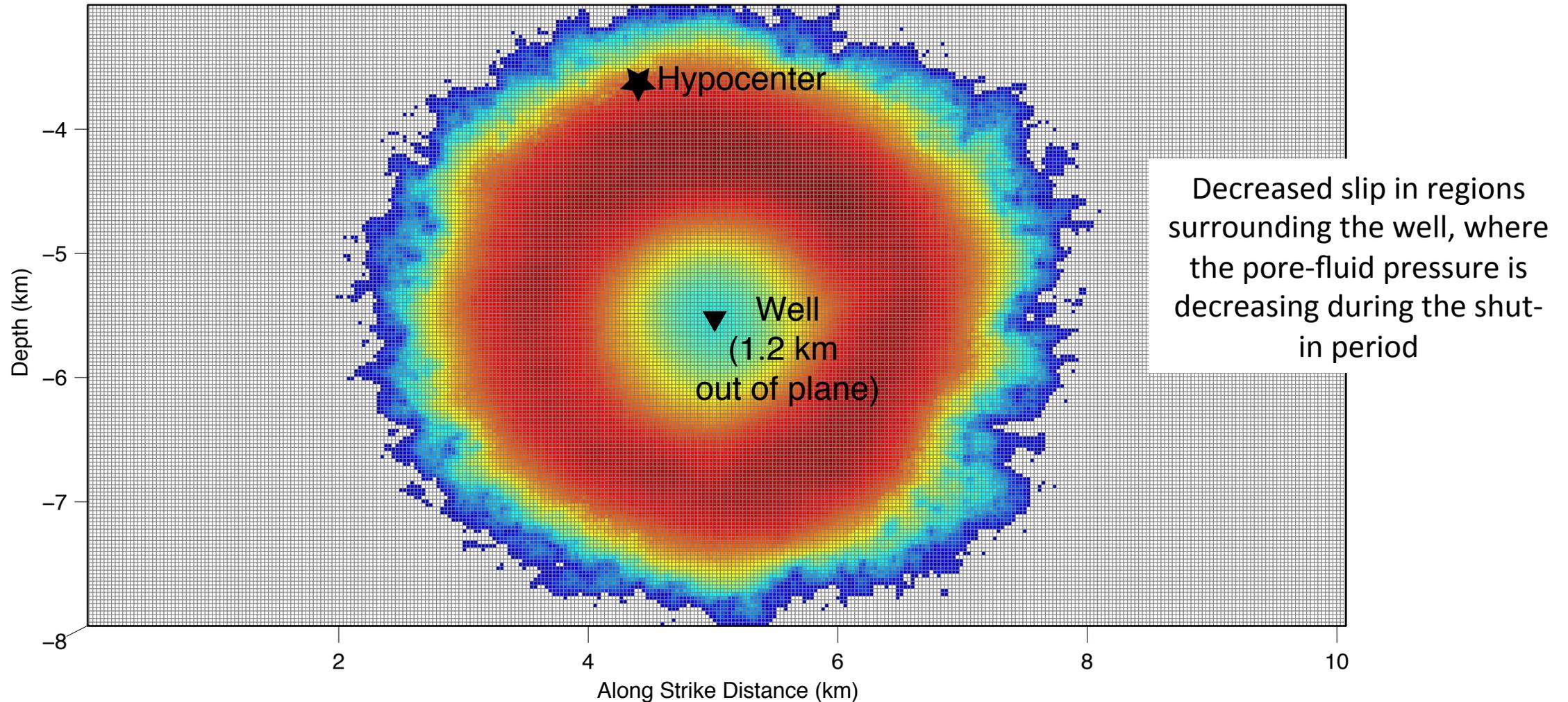
Overpressure on hypocentral element is rising at the time of the M4.5



Results: Earthquake Magnitude with Time (Periodic Injection Rate)

Event # 3986; M = 4.8; dt- = -0.09 secs; dt = 0; dt+ = -0.09 secs

Origin time (yrs): 1032.688 Nucleated on patch 6532 (NA) max slip = 0.041 m full color scale slip = 0.041 m



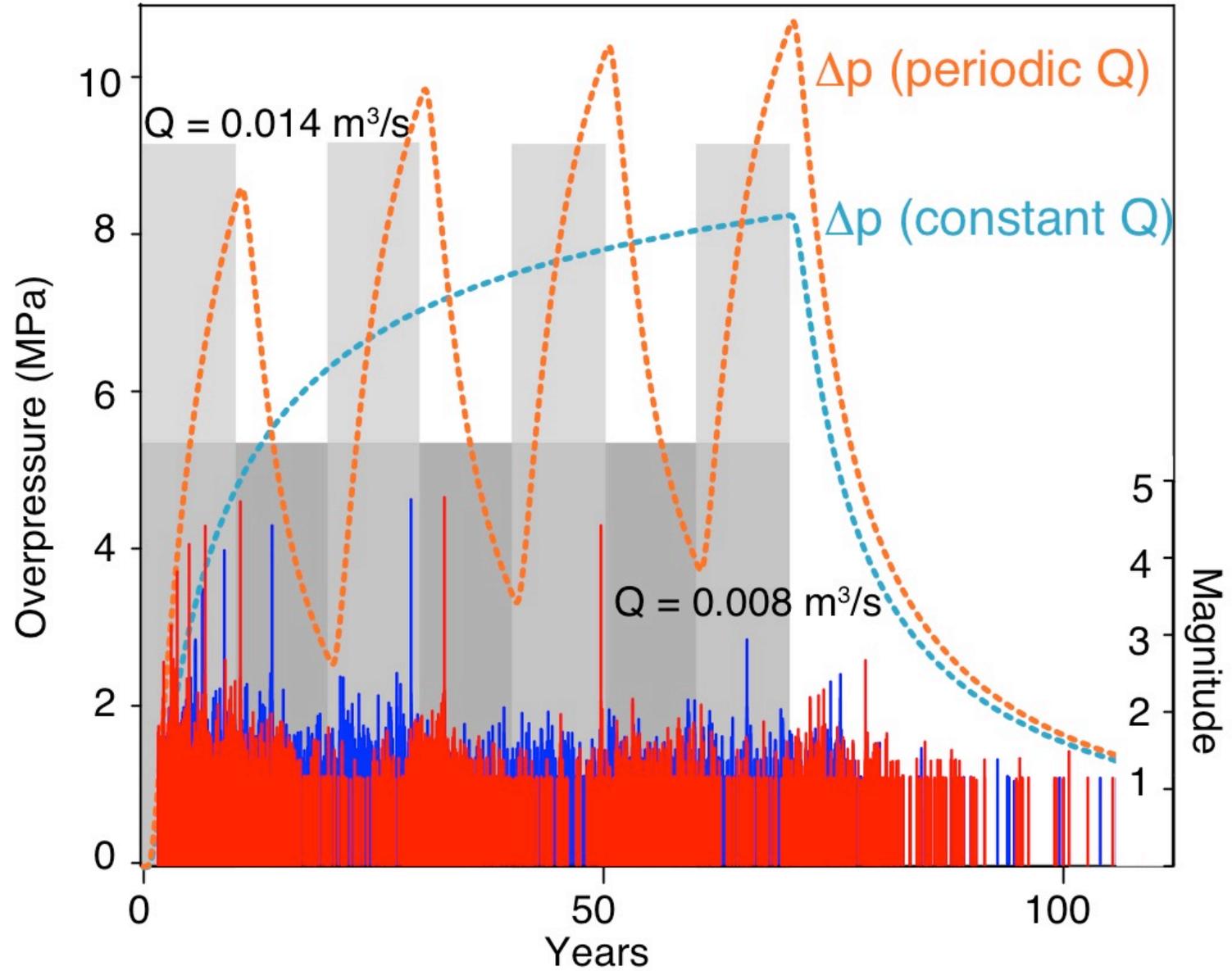
Results: Compare Earthquake Magnitudes with Time

Constant Q:

- Fewer large events (4 $M > 3.5$)
- Earliest large events are smaller and have longer inter-event times

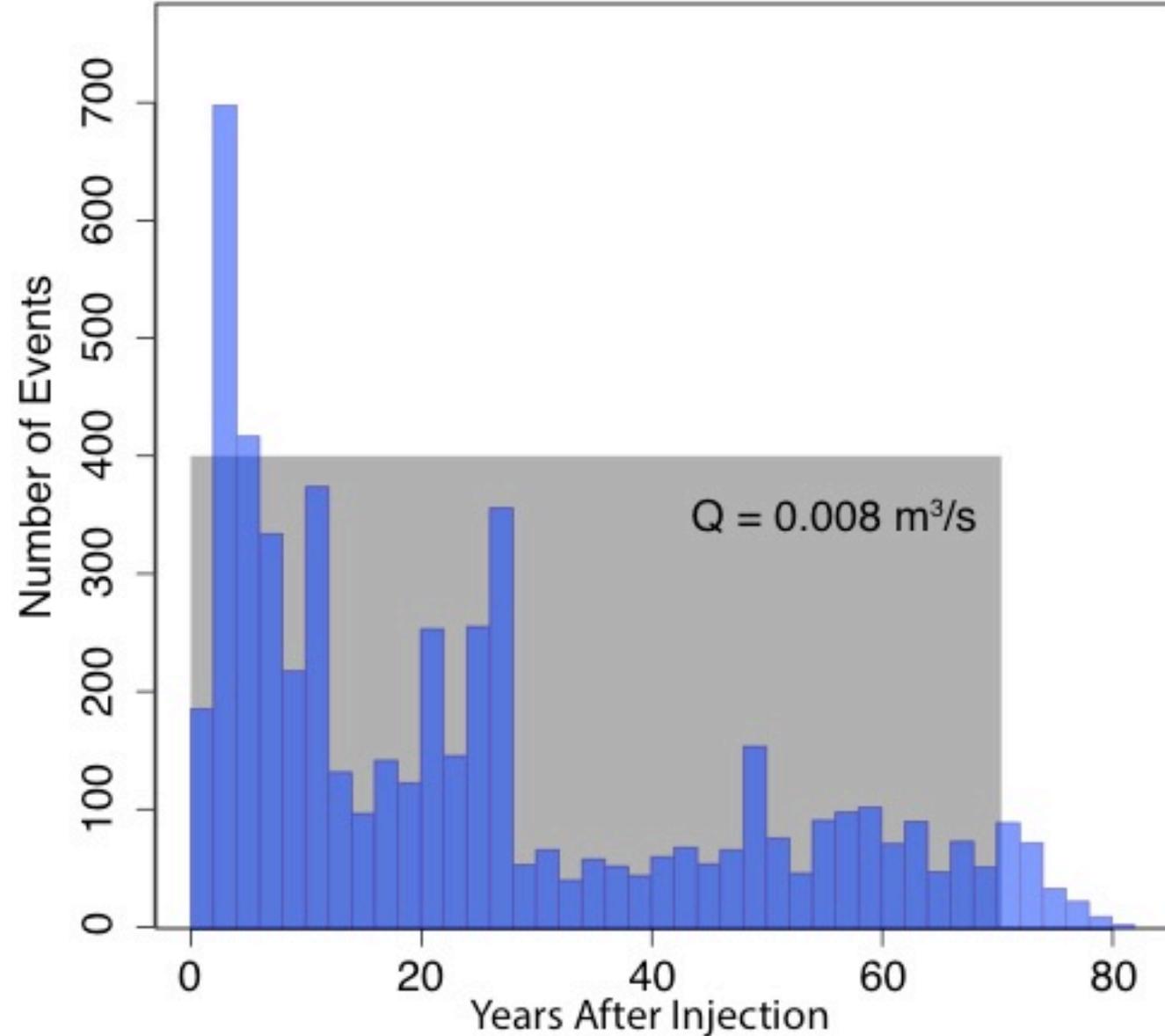
Periodic Q:

- More large events (7 $M > 3.5$), especially early in the sequence
- Earliest large events are bigger and have shorter inter-event times
- Longer inter-event time after first injection period



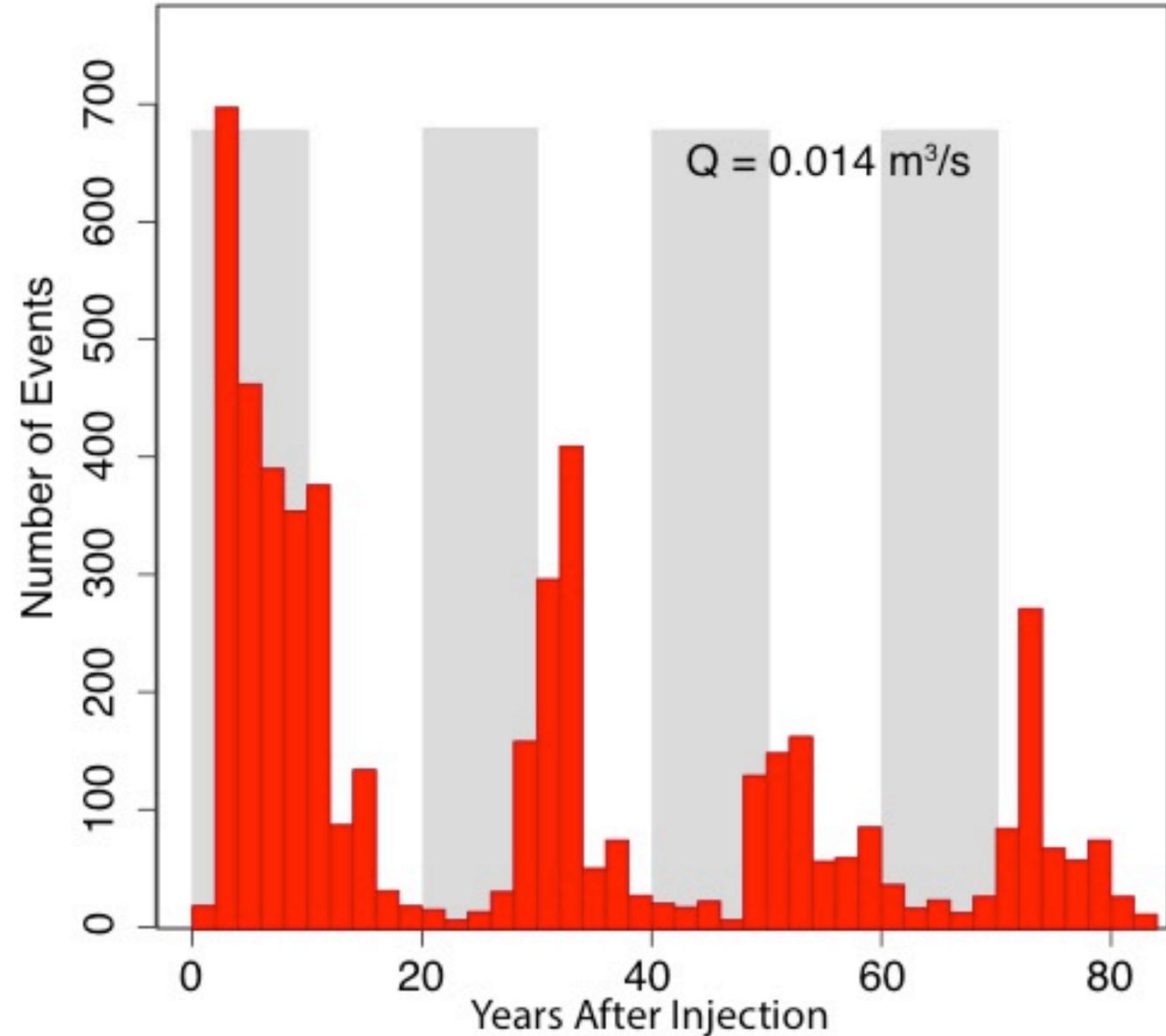
Results: Number of Events with Time (Constant Q)

- Most events occur shortly following injection
- Event rate decays with time



Results: Number of Events with Time (Periodic Q)

- Early events occur mostly during first injection period
- Later events primarily occur during the time shut-in cycle
 - Aftershocks following post shut-in large events
- Event rate fluctuates with time, but decreases overall



Results: Compare Number of Events with Time

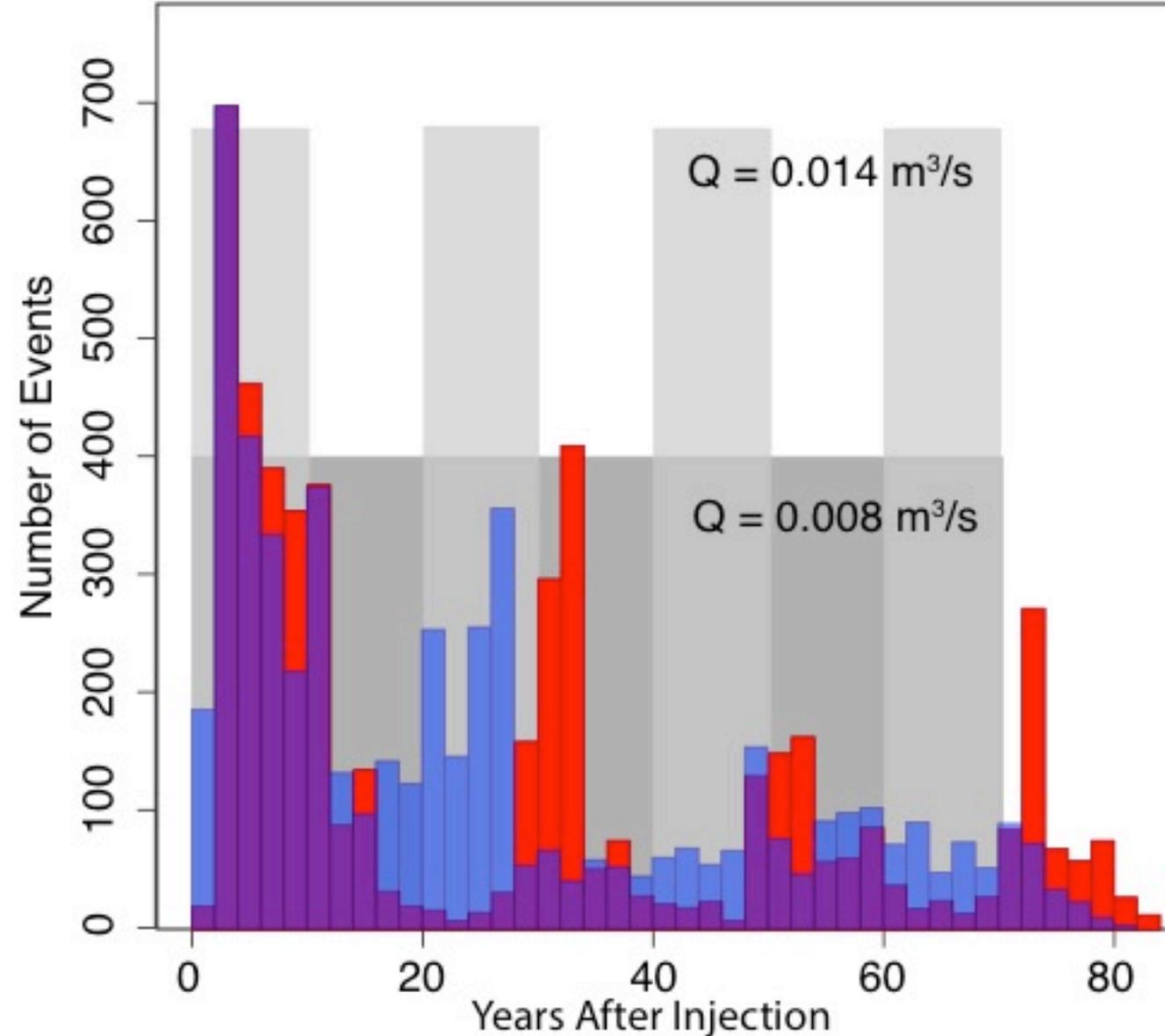
- Relatively similar history during the first 15 years

Constant Q:

- Event rates decay with time
- Highest rates within the first 25 years

Periodic Q:

- Highly variable earthquake rate
- Several periods of increased rate



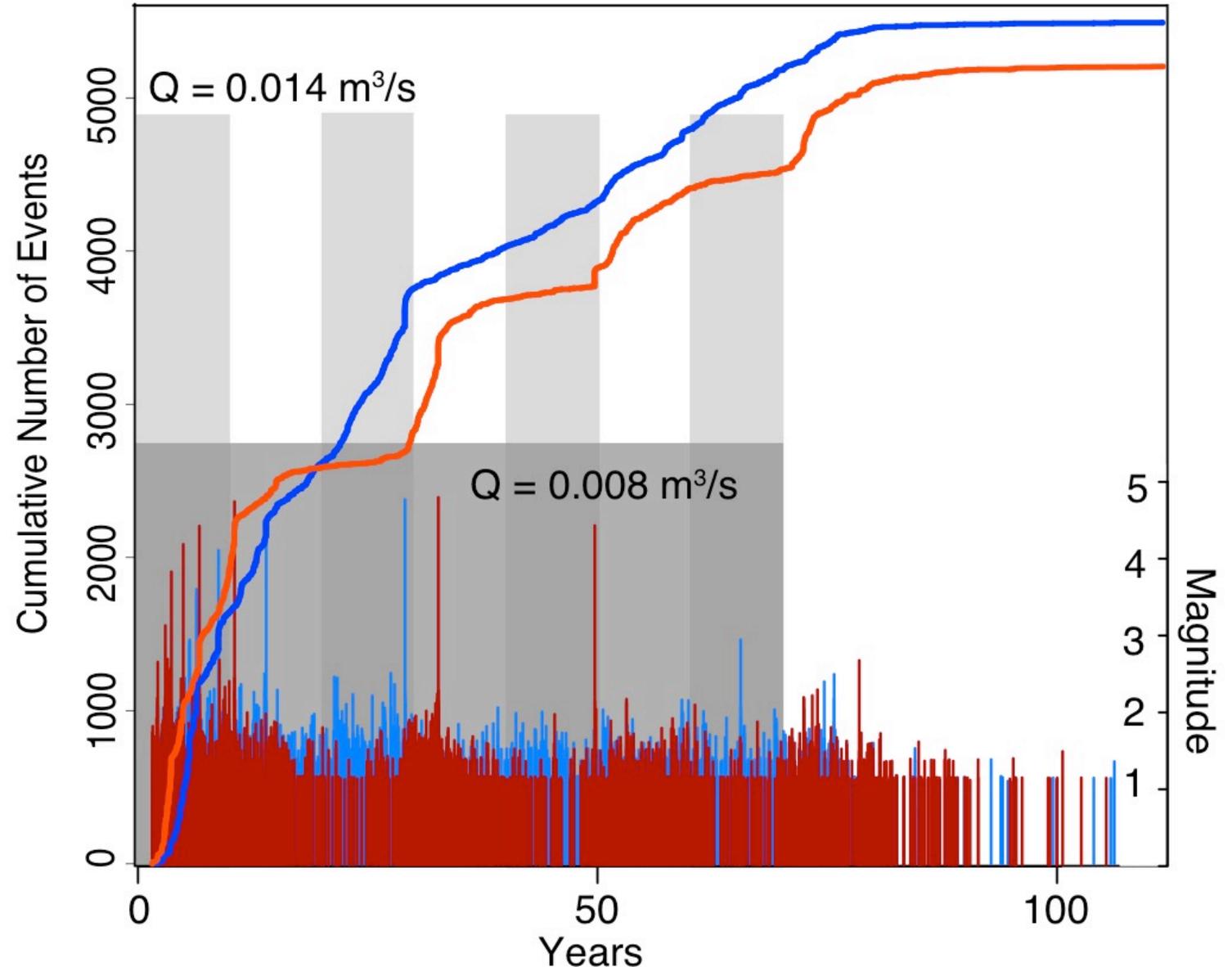
Results: Compare Cumulative Number of Events with Time

Constant Q:

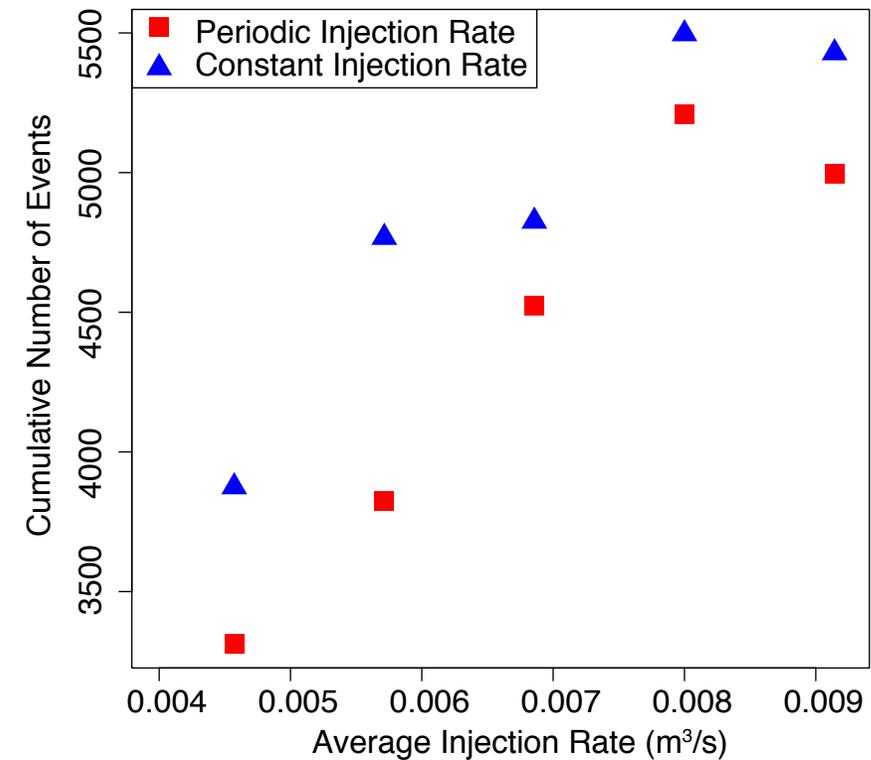
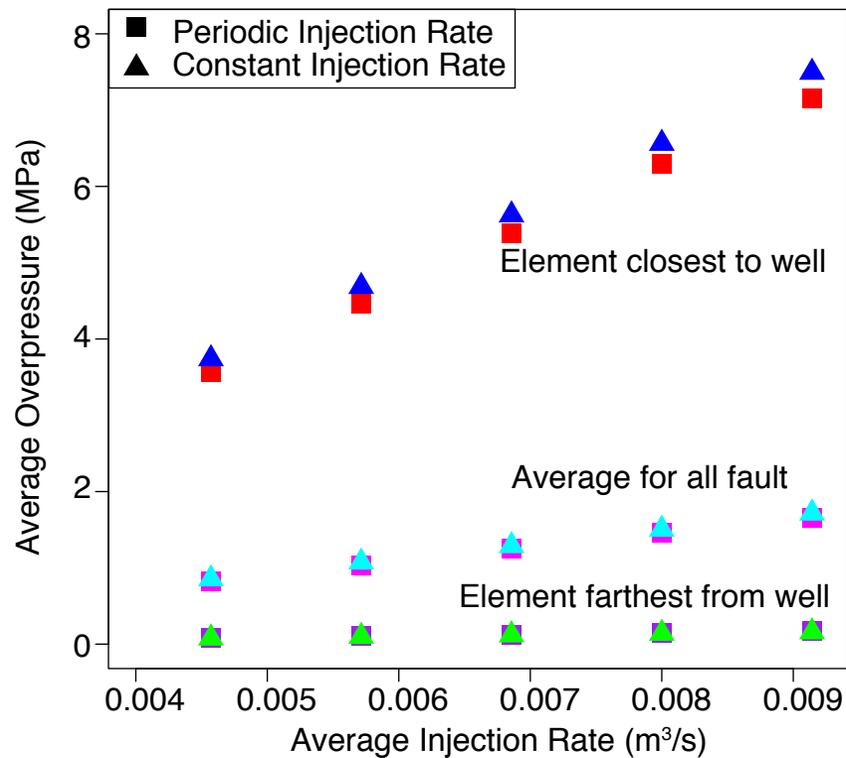
- Three distinct periods of events:
 - 1000 – 1025: highest rate
 - 1025 – 1075: moderate rate
 - 1075+: quiescence
- More events total

Periodic Q:

- Highly variable earthquake rate
- Apparent increased rate leading up to large events
- Fewer events total

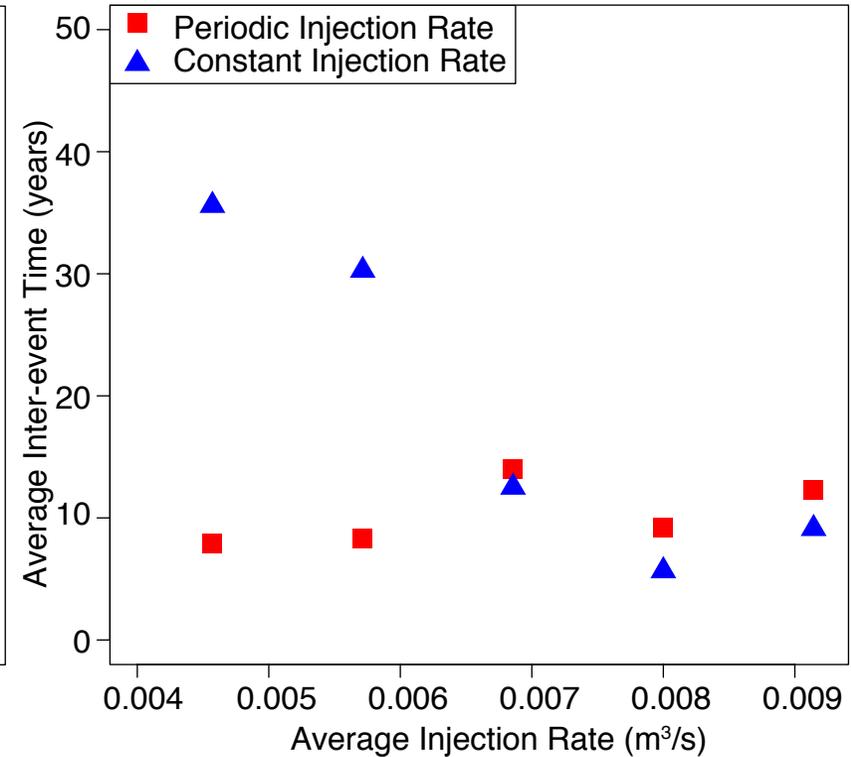
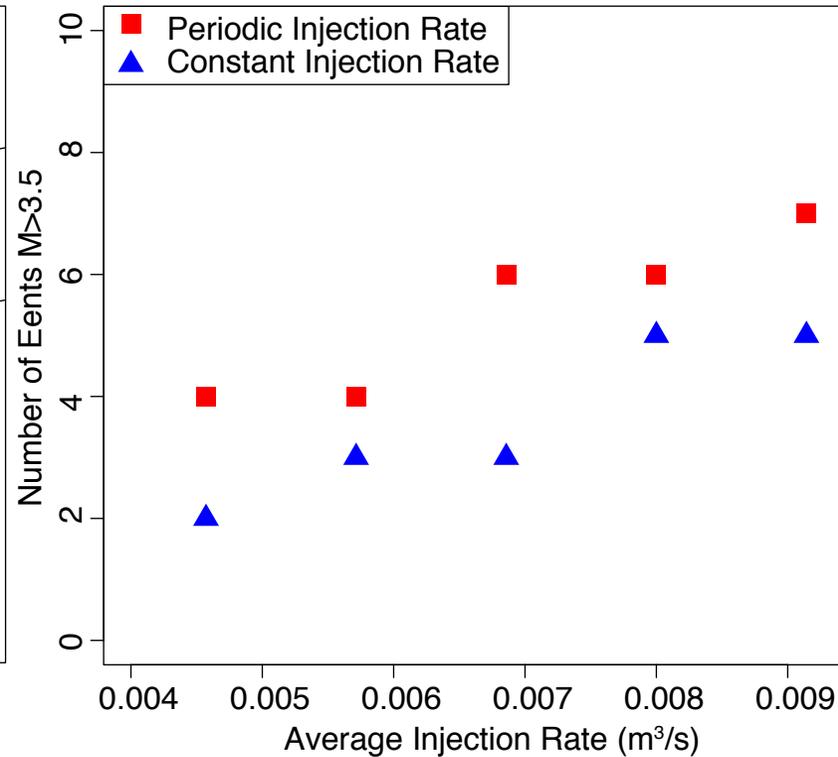
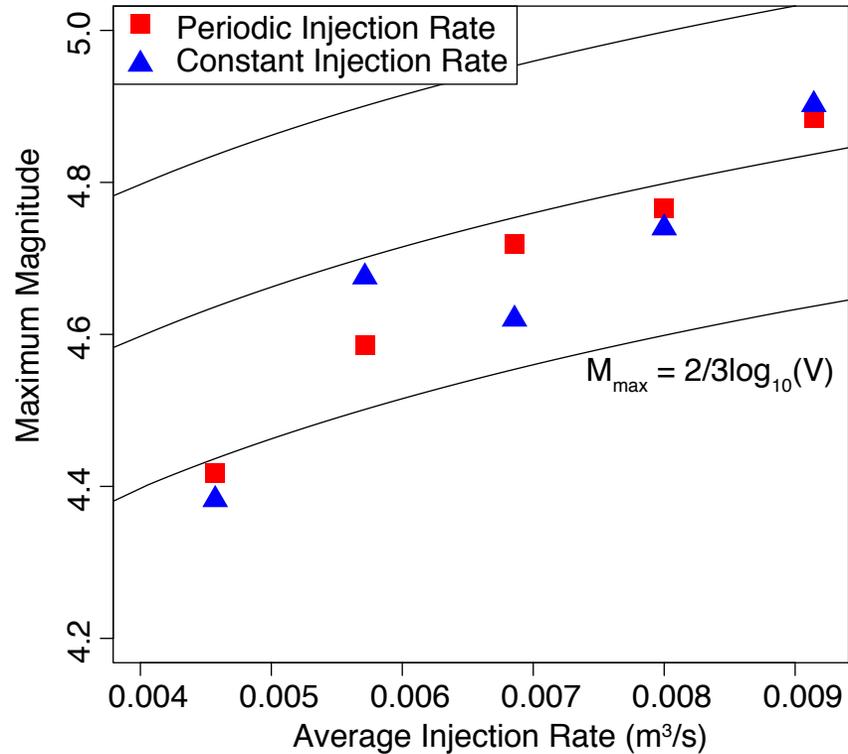


Summary: Constant vs. Periodic Injection Rates



- Overpressure is greater for period injection
- Average overpressure increases with increasing injection rate
- Constant injection rates lead to large total number of events

Summary: Constant vs. Periodic Injection Rates



- M_{\max} increases with increased injection rate
- M_{\max} is not systematically larger by injection rate.
- Slope is larger than M_{\max} scaling by injected volume.

• More events $M > 3.5$ with periodic injection history

- Inter-event times are inversely related in constant average injection rates
- Relatively unaffected by periodic injection rates

Conclusions

- Rate-state constitutive parameters highly influence induced sequences

Injection Methods (with roughly equivalent average overpressures over time):

1. Low constant injection rates
 - More events overall
 - Fewer large events
 - Larger events have longer inter-event times
2. Periodically injecting at high rates
 - Fewer events
 - More large events
 - Shorter inter-event times

Acknowledgements

Thank you!

- Elizabeth S. Cochran
- NSF
- SCEC

Questions?

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